

Biogenic Contributions to Methane Trends from 1990 to 2004

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1. Introduction

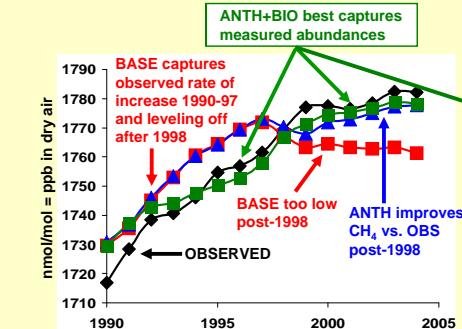
Methane (CH_4) emission controls are a cost-effective strategy for abating global surface ozone, while simultaneously slowing greenhouse warming [West and Fiore, 2005]. Sources of CH_4 are shown to the right; the major CH_4 sink is reaction with the hydroxyl radical (OH) in the troposphere. Surface CH_4 concentrations rose by 5–6 ppb yr^{-1} on average until 1999 when they leveled off (see Figure below). There is no clear consensus on the driving mechanism.

Possible explanations suggested by previous studies:

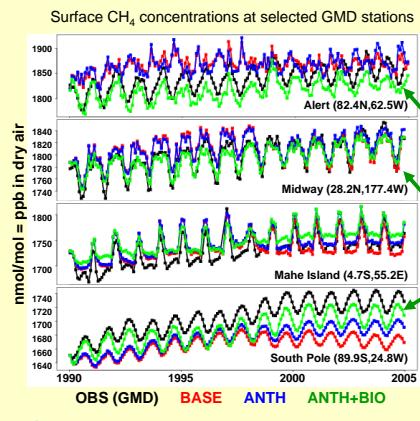
- (1) Source changes of CH_4 [e.g. Langenfelds et al., 2002; Wang et al., 2004] or other species [e.g. Karlsson and Isaksen, 2000]
- (2) Meteorologically-driven changes in the CH_4 sink [e.g. Warwick et al., 2002; Dentener et al., 2003; Wang et al., 2004]
- (3) Approach to steady-state with constant lifetime [Dlugokencky et al., 2003]

What is the contribution from each process?
Are existing bottom-up CH_4 emission inventories for the 1990s consistent with observations?

3. Influence of Sources on Surface CH_4 Distribution and Trend



Global mean surface CH_4 concentrations as measured (or sampled in the model) at 42 Global Monitoring Division (GMD) stations [e.g. Dlugokencky et al., 2005] with an 8-year minimum record. Values are area-weighted after averaging in latitudinal bands (60–90°N, 30–60°N, 0–30°S, 30–90°S).



REFERENCES

- Dentener, F., et al. (2003), *J. Geophys. Res.*, 108, 4424, doi:10.1029/2002JD002916.
Karlsdóttir, S., and I.S.A. Isaksen (2000), *Geophys. Res. Lett.*, 27 (1), 93–96.
Langenfelds, R.L., et al. (2002), *Global Biogeochem. Cycles*, 16, 1048, doi:10.1029/2001GB001466.
Oliver, J.G.J., et al. (1999), *Environmental Science & Policy*, 2, 241–264.
Olivier, J.G.J. (2002) In: "CO₂ emissions from fuel combustion 1971–2000", 2002 Edition, pp. III–III-31. International Energy Agency (IEA), Paris. ISBN 92-64-09794-5.

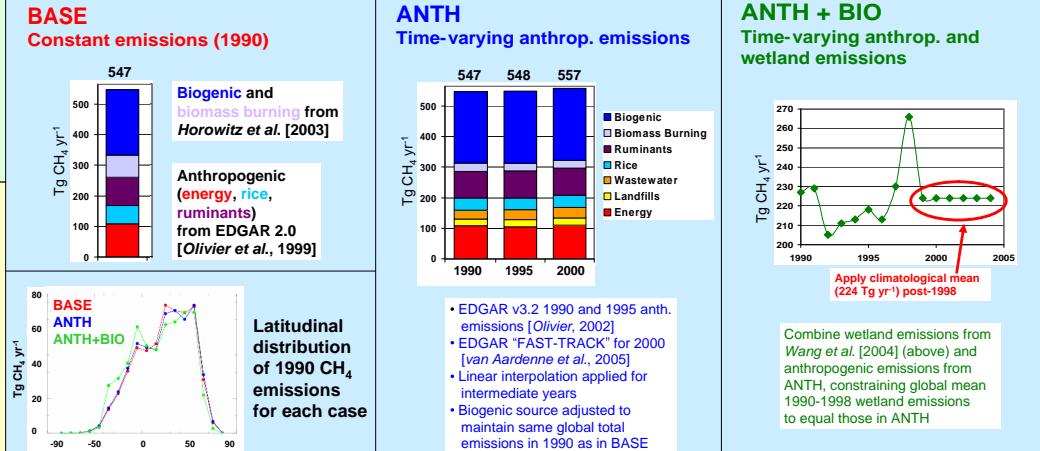
Dlugokencky, E.J., et al. (2005), *J. Geophys. Res.*, 110, D18306, doi:10.1029/2005JD006035.

Horowitz, L.W., et al. (2003), *J. Geophys. Res.*, 108, 4784, doi:10.1029/2002JD002853.



2. Methane in the MOZART-2 CTM

Sensitivity simulations applying different CH_4 emission inventories:

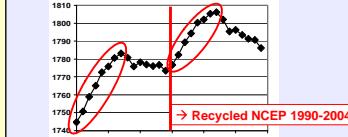


- EDGAR v3.2 1990 and 1995 anth. emissions [Olivier, 2002]
- EDGAR "FAST-TRACK" for 2000 [van Aardenne et al., 2005]
- Linear interpolation applied for intermediate years
- Biogenic source adjusted to maintain same global total emissions in 1990 as in BASE

- ~100 gas and aerosol species, ~200 reactions
- NCEP meteorology 1990–2004
- 1.9° latitude x 1.9° longitude x 64 vertical levels
- detailed description in Horowitz et al. [2003]

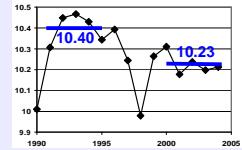
4. Meteorologically-driven Changes in the CH_4 Lifetime

Area-weighted global mean surface CH_4 in BASE simulation (constant emissions)



→ Meteorological drivers for observed trend
→ Not just simple approach to steady-state

CH_4 Lifetime Against Tropospheric OH (τ)



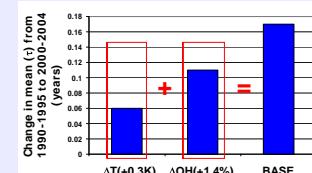
→ Mean annual CH_4 lifetime shortens

How might meteorology affect τ ?

$$\tau = \frac{[CH_4]}{k[OH][CH_4]}$$

Transport to sink regions
Temperature
Humidity
Lightning NO_x
Photolysis (changes in stratospheric O₃ columns neglected)
Changes in emissions of other species (neglected)

Deconstruct $\Delta\tau$ (-0.17 years) from 1991–1995 to 2000–2004 into individual contributions by varying temperature and OH separately



→ OH increases in the model by +1.4% from 1991–1995 to 2000–2004 due to a 0.3 Tg yr^{-1} (~10%) increase in lightning NO_x

→ Large uncertainties in magnitude (and trend) of lightning NO_x emissions in the real atmosphere

5. Conclusions

- BASE simulation (constant emissions) captures observed rate of CH_4 increase from 1990–1997, and leveling off post-1998
- ANTH emissions (EDGAR 3.2 + FT-2000 inventories) improve modeled CH_4 post-1998
- ANTH+BIO spatial and temporal distribution of wetland CH_4 emissions is critical for reproducing the observed seasonality, interhemispheric gradient, and global mean trend
- τ_{CH_4} decreases by ~2% from 91–95 to 00–04 due to warmer temperatures (35%) and higher OH (65%, resulting from a ~10% increase in lightning NO_x emissions)

Future research should:

- consider climate-driven feedbacks from biogenic emissions and fires (CH_4 , CO, NO_x, NMVOC) on τ_{CH_4}
- develop more physically-based parameterizations of lightning NO_x emissions to determine whether higher emissions are a robust feature of a warmer climate